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SPECIAL

Modern seawater acidification: the response of foraminifera to high-CO₂ conditions in the Mediterranean Sea

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The seas around the island of Ischia (Italy) have a lowered pH as a result of volcanic gas vents that emit carbon dioxide from the sea floor at ambient seawater temperatures. These areas of acidified seawater provide natural laboratories in which to study the long-term biological response to rising CO₂ levels. Benthic foraminifera (single-celled protists) are particularly interesting as they have short life histories, are environmentally sensitive and have an excellent fossil record. Here, we examine changes in foraminiferal assemblages along pH gradients at CO₂ vents on the coast of Ischia and show that the foraminiferal distribution, diversity and nature of the fauna change markedly in the living assemblages as pH decreases.

Short-term laboratory experiments indicate that, with rising atmospheric *p*CO₂ and a lowering of oceanic pH, many organisms with calcareous shells may be unable to survive the next 100 years (Orr *et al.* 2005; Doney *et al.* 2009; Moy *et al.* 2009; Veron *et al.* 2009) although others appear to be resistant (Findlay *et al.* 2009; Ries *et al.* 2009). A wide range of benthic macro-organisms (corals, barnacles, gastropods, echinoderms, sea-grasses, calcareous and non-calcareous algae) living in the shallow-water environments around the coast of Ischia (Bay of Naples, Italy) show marked responses to increasing levels of CO₂ in the water (Hall-Spencer *et al.* 2008; Martin *et al.* 2008; Rodolfo-Metalpa *et al.* 2010). Around Castello Aragonese (Fig. 1) shallow-water habitats have zones with increasing CO₂ levels where mean pH values (measured in Total Scale at the sediment–water interface) range from 8.14 to 6.57 (for details, see Hall-Spencer *et al.* 2008). These waters provide a location at which predicted effects of ocean acidification (Caldeira & Wickett 2005; Royal Society 2005; IPCC 2007; Feely *et al.* 2004) can be tested by a study of the distribution of the living biota. The seawater adjacent to Ischia is acidified by gas

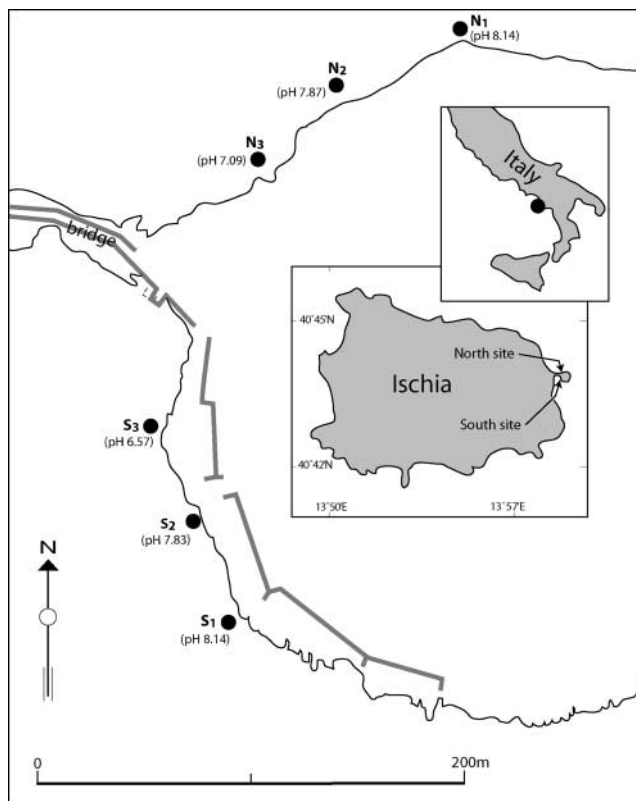


Fig. 1. Location of Ischia and sampling locations on the Castello Aragonese. Samples were collected from sites on traverses undertaken both north and south of the bridge and the foraminiferal assemblages were compared.

comprising 90.1–95.3% CO₂, 3.2–6.6% N₂, 0.6–0.8% O₂, 0.08–0.1% Ar and 0.2–0.8% CH₄ (no sulphur). The salinity of the water (38‰) and total alkalinity (2.5 mequiv. kg⁻¹) are uniform between all the sampling stations and the annual temperature range is 13–25 °C (Hall-Spencer *et al.* 2008).

Samples and methods. We have studied the assemblages of foraminifera from samples collected in the shallow waters around Ischia. We used a trowel to collect 1 cm deep 200 ml surface sediment samples at 2 m depth from areas with mean ambient pH levels of 8.14, 7.87, 7.83, 7.09 and 6.57. The sediment samples were collected by diving, and they were stored in buffered formalin, washed on a 63 μm sieve, stained with Rose Bengal (to distinguish live from dead individuals), re-washed and dried at <40 °C. Normal micropalaeontological dry sieving and counting techniques were employed. The foraminifera were identified (in the first instance using Cimerman & Langer 1991) and imaged using a JEOL 5600 scanning electron microscope with a digital capture system.

The living assemblages from the sea-floor areas with a ‘normal’ pH of *c.* 8.14 contain a range of taxa dominated by miliolids, as would be expected in a shallow-water Mediterranean environment with a salinity of 38‰ (Murray 2006). The dominant taxa are *Elphidium* sp. cf. *E. advenum*, *Triloculina tricarinata*, *Pyrgo* sp., *Miliolinella elongata* and *Peneroplis planatus* (Fig. 2). Samples (S2 and N2 in Fig. 1) with lowered pH (7.87–7.83) contain *Miliammina fusca*, *Trochammina inflata*, *Ammonia tepida* and *Reophax* sp. Samples (S3 and N3 in Fig. 1)

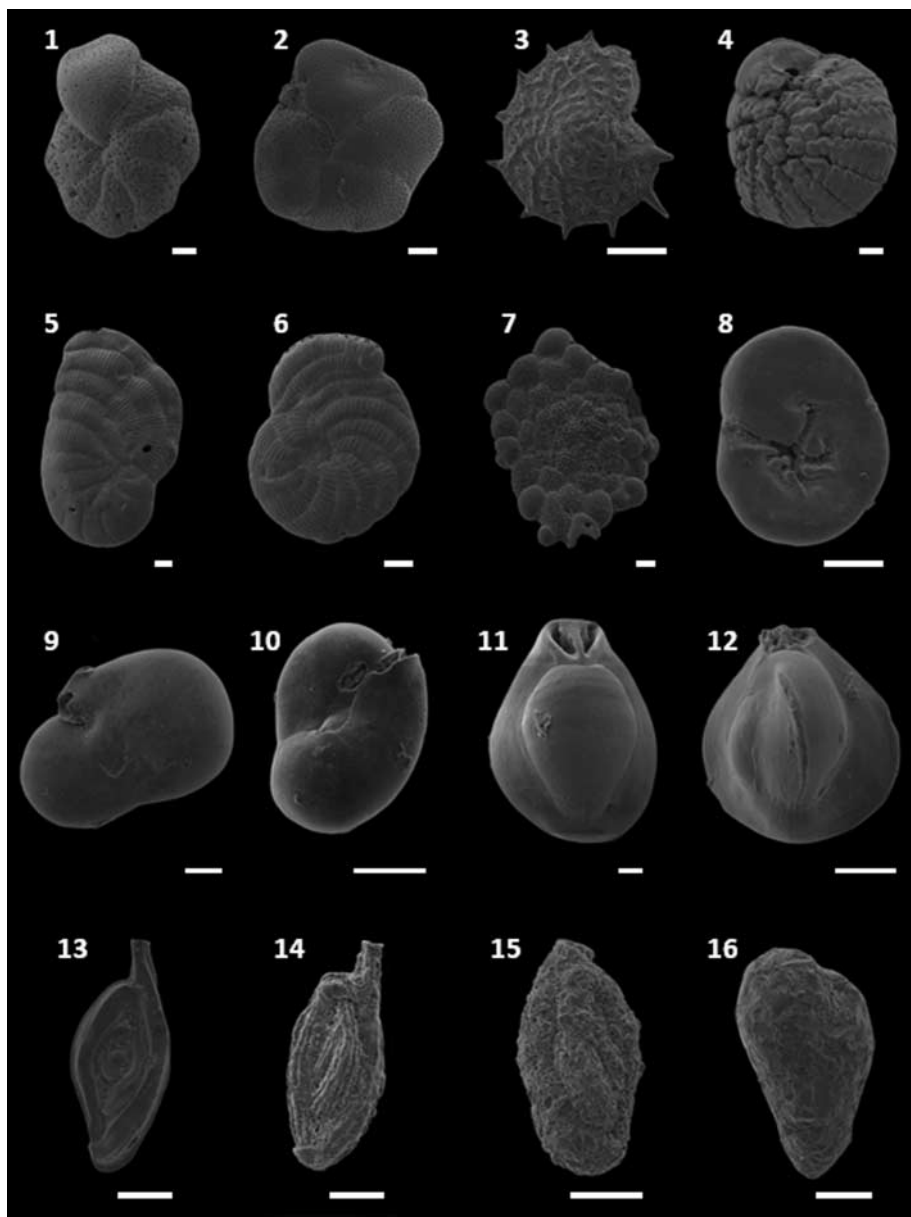


Fig. 2. Selected species from Ischia, Italy. 1, *Lobatula lobatula*; 2, *Cibicides refulgens*; 3, *Elphidium aculeatum*; 4, *Ammonia tepida*; 5, *Peneroplis planatus*; 6, *Laevipeneroplis karreri*; 7, *Planorbulina mediterraneensis*; 8, *Rosalina bradyi*; 9, *Parrina bradyi*; 10, *Miliolinella labiosa*; 11, *Quinqueloculina seminula*; 12, *Triloculina tricarinata*; 13, *Spiroloculina ornata*; 14, *Sigmoininita costata*; 15, *Miliammina fusca*; 16, *Textularia* sp. cf. *T. bocki*. Scale bar represents 100 μm . It is important to note that, in the case of specimens collected in the areas of normal pH, the preservation is excellent.

from areas with the highest CO_2 levels (mean pH 7.09–6.57) contain (at a pH of 7.6) an impoverished assemblage of agglutinated foraminifera that includes *Trochammina inflata*, *Miliammina fusca*, *Textularia* sp. cf. *T. bocki* and *Ammoglobigerina globigeriniformis*.

Impact of acidification. The reduction of the foraminiferal diversity (and abundance) mirrors that recorded for macrobenthic calcifying organisms (Hall-Spencer *et al.* 2008; Martin *et al.* 2008). Figure 3 shows that diversity falls from 24 species to four species from the normal marine samples towards the areas with high CO_2 levels. The foraminiferal assemblage also shifts from one dominated by calcareous forms to one dominated by agglutinated taxa within a distance of <200 m as a result of the acidification of the water by CO_2 with no other measurable gradients; depth, salinity, temperature, granulometry and light levels were similar across all sites (Hall-Spencer *et al.* 2008). The change in the assemblages is illustrated by use of the

triangular diagram developed by Murray for the investigation of foraminiferal ecology (Murray 1991a, pp. 232–244; Murray 1991b). The samples from Ischia with normal pH are seen to plot in the hypersaline field (dominated by miliolids), but with reducing pH the data points migrate towards the sector with 100% agglutinated (textulariid) taxa (Fig. 4).

Implications for the fossil record. Samples with c. 100% agglutinated assemblages of foraminifera are well known from the geological record (Scott *et al.* 1983; Jones 1988; Charnock & Jones 1990; Czarniecki 1993) and there has been much debate as to whether these are primary (a response to the environment) or secondary (the result of dissolution during taphonomy or diagenesis) (see Hart 1983, pp. 251–263, fig. 3). In the 1990s Alve and Murray conducted a series of acidification experiments to test the fate of a ‘normal’ foraminiferal assemblage when subjected to treatment with weak acid (Murray 1989, 2006; Alve & Murray 1994, 1995; Murray & Alve 1994, 1999a,b, 2000). The assem-

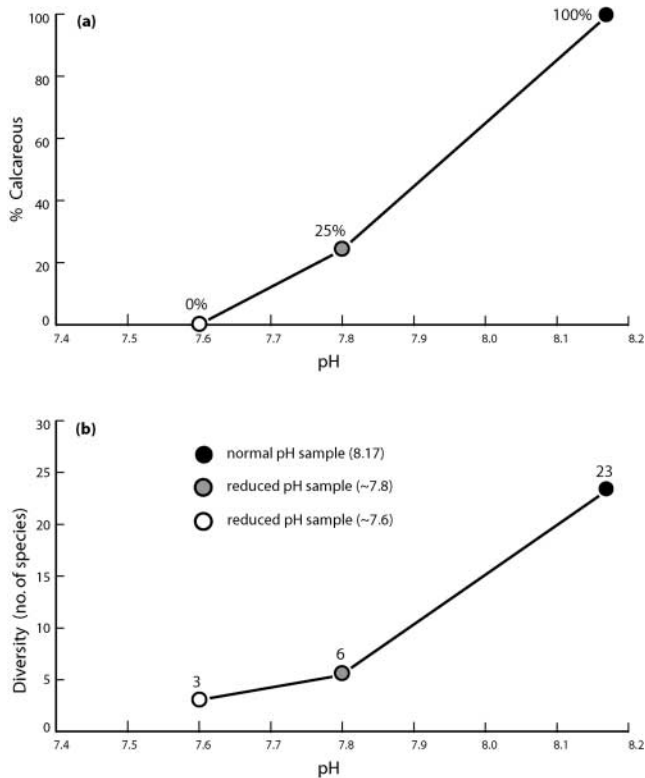


Fig. 3. (a) Graph of changing composition of the living assemblage (calcareous foraminifera:agglutinated foraminifera) across the range of pH values: (b) diversity (measured as total number of living species) across the range of pH values.

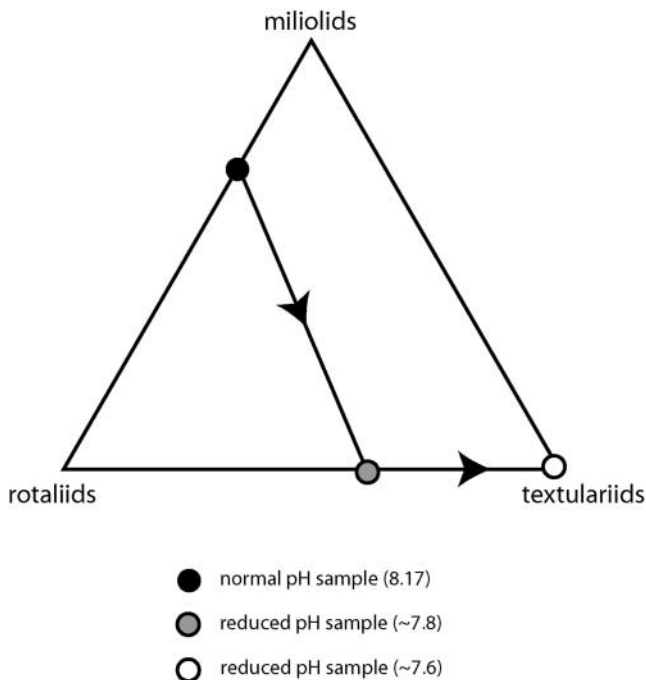


Fig. 4. Triangular diagram plotting miliolid taxa, agglutinated taxa and rotaliid taxa, which shows the changing composition of the living assemblage with the recorded changes in pH. It should be noted that the change in pH required to completely change the nature of the assemblage is relatively significant (from pH 8.2 to pH 7.8) although intermediate values have not yet been investigated.

blages generated (acid-treated assemblages or ATAs) by Alve & Murray (1995) were, not surprisingly, composed of 100% agglutinated taxa lacking calcareous cement. In the case of the samples from Ischia, the change from a calcareous-dominated assemblage (75%) to one with no calcareous taxa appears to be entirely due to changing pH and not post-death acidification during burial and diagenesis. In another example of the dominance of agglutinated taxa in an assemblage (Murray *et al.* 2003) there was a restriction caused by depleted levels of dissolved oxygen in the water column (though not anoxia) and pH was not recorded.

The Palaeocene–Eocene Thermal Maximum (PETM) is one of the most significant events in the Cenozoic and there is a body of evidence to suggest that there was a period of ocean acidification, which, in turn, affected the marine isotope record that is based on benthic foraminifera (Spero *et al.* 1997; Zachos *et al.* 2005; Pagani *et al.* 2006). The data from Ischia certainly confirm that a change in pH could cause a marked change in the benthic assemblage and even ‘extinction’ from a geological succession (Speijer *et al.* 1996; Zachos *et al.* 2005; Kaiho *et al.* 2006; Giusberti *et al.* 2009). Nguyen *et al.* (2009) have recently reported on dissolution experiments (using water with a pH value as low as 6.6) aimed at understanding the events during the PETM. The data from Ischia, however, provide information on the pH conditions under which the living assemblage is affected, rather than providing information on post-mortem dissolution. Presence or absence of foraminifera at the PETM must be a combination of both a biological restriction and a diagenetic impact.

Summary. The living (and dead) foraminiferal assemblages showed significant changes resulting from CO₂ acidification of the shallow waters around Ischia. The assemblages in normal conditions (pH 8.2–8.14) were dominated by calcareous forms including abundant miliolids (*Triloculina*, *Pyrgo*, *Miliolinella*, *Quinqueloculina* and *Peneroplis*). The most impoverished assemblages (pH 7.6) were all agglutinated species of *Trochammina*, *Miliammina*, *Textularia* and *Ammoglobigerina*. This work confirms that, in open-water marine conditions, foraminifera are sensitive to the effects of ocean acidification, thereby indicating that rising levels of atmospheric pCO₂ may cause significant changes in these assemblages. It also confirms that changes in pH could have caused the extinctions recorded at events such as the PETM.

Appendix. Foraminifera recorded in the samples from Ischia. The original author names and dates are given (for references, see Cimerman & Langer 1991).

Normal pH 8.2–8.14

Ammonia inflata (Seguenza 1862); *Brizalina* sp.; *Cibicides advenum* (d’Orbigny 1839); *Cibicides refulgens* Montfort 1808; *Cycloforina tenuicollis* (Wiesner 1923); *Elphidium aculeatum* (d’Orbigny 1846); *Elphidium* sp. cf. *E. advenum* (Cushman 1922); *Lepidodeuterammina* sp.; *Massilina gualtieriana* (d’Orbigny 1839); *Miliolinella elongata* Kruit 1955; *Miliolinella labiosa* (d’Orbigny 1839); *Parrina bradyi* (Millett 1898); *Peneroplis pertusus* (Forskål 1775); *Peneroplis planatus* (Fichtel & Moll 1798); *Planorbulina mediterraneanensis* d’Orbigny 1826; *Pyrgo* sp.; *Quinqueloculina berhelotiana* d’Orbigny 1839; *Quinqueloculina jugosa* Cushman 1944; *Quinqueloculina seminula* (Linné 1758); *Rosalina vilardeboana* d’Orbigny 1839; *Triloculina tricarinata* d’Orbigny 1826; *Triloculina* sp.

Medium pH 7.8

Ammonia tepida (Cushman 1926); *Miliammina fusca* (Brady 1870); *Reophax* sp.; *Trochammina inflata* (Montagu 1803).

Low pH 7.6

Ammoglobigerina globigeriniformis (Parker & Jones 1865); *Miliammina fusca* (Brady 1870); *Textularia* sp. cf. *T. bocki* Höglund 1947; *Trochammina inflata* (Montagu 1803).

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References

- ALVE, E. & MURRAY, J.W. 1994. Ecology and taphonomy of benthic foraminifera in a temperate mesotidal inlet. *Journal of Foraminiferal Research*, **24**, 18–27.
- ALVE, E. & MURRAY, J.W. 1995. Experiments to determine the origin and palaeoenvironmental significance of agglutinated foraminiferal assemblages. In: KAMINSKI, M.A., GEROGH, S. & GASIŃSKI, M.A. (eds) *Proceedings of the Fourth International Workshop on Agglutinated Foraminifera, Kraków, Poland, September 12–19, 1993*. Grzybowski Foundation Special Publication, **3**, 1–11.
- CALDEIRA, K. & WICKETT, M.E. 2005. Ocean model predictions of chemistry changes from carbon dioxide emissions to the atmosphere and ocean. *Journal of Geophysical Research*, **110**, C09504, doi:10.1029/2004JC002671.
- CHARNOCK, M.A. & JONES, E.W. 1990. Agglutinated foraminifera from the Palaeogene of the North Sea. In: HEMLEBEN, C., KAMINSKI, M.A., KUHN, W. & SCOTT, D.B. (eds) *Paleoecology, Biostratigraphy and Taxonomy of Agglutinated Foraminifera*. NATO, ASI Series, **C-327**, 139–244.
- CIMERMAN, F. & LANGER, M.R. 1991. *Mediterranean Foraminifera*. Slovenska akademija znanosti in umetnosti, Ljubljana.
- CZARNIECKI, S. 1993. Grzybowski and his school: the beginnings of applied micropalaeontology in Poland at the turn of the 19th and 20th centuries. In: KAMINSKI, M.A., GEROGH, S. & KAMINSKI, D.G. (eds) *The Origins of Applied Micropalaeontology: The School of Józef Grzybowski*. Grzybowski Foundation Special Publication, **1**, 1–15.
- DONEY, S.C., FABRY, V.J., FEELY, R.A. & KLEYPAS, J.A. 2009. Ocean acidification: the other CO₂ problem. *Annual Review of Marine Science*, **1**, 169–192, doi:10.1146/annurev.marine.010908.163834.
- FEELY, R.A., SABINE, C.L., LEE, K., BERELSON, W., KLEYPAS, J., FABRY, V.J. & MILLERO, F.J. 2004. Impact of anthropogenic CO₂ on the CaCO₃ system in the oceans. *Science*, **305**, 362–366, doi:10.1126/science.1097329.
- FINDLAY, H.S., WOOD, H.L., KENDALL, M.A., SPICER, J.I., TWITCHETT, R.J. & WIDDICOMBE, S. 2009. Calcification, a physiological process to be considered in the context of the whole organism. *Biogeosciences Discussions*, **6**, 2267–2284.
- GIUSBERTI, L., COCCIONI, R., SPROVERI, M. & TATEO, F. 2009. Perturbation at the sea floor during the Paleocene–Eocene Thermal Maximum: evidence from the benthic foraminifera at Contessa Road, Italy. *Marine Micropaleontology*, **70**, 102–119, doi:10.1016/j.marmicro.2008.11.003.
- HALL-SPENCER, J.M., RODOLFO-METALPA, R., MARTIN, S., ET AL. 2008. Volcanic carbon dioxide vents show ecosystem effects of ocean acidification. *Nature*, **454**, 96–99, doi:10.1038/nature07051.
- HART, M.B. 1983. Some thoughts on the ecology (and palaeoecology) of the arenaceous foraminifera: a workshop report. In: VERDENIUS, J.G., VAN HINTE, J.E. & FORTUIN, A.R. (eds) *Proceedings of the First Workshop on Arenaceous Foraminifera, 7–9 September 1981*. Institutt for Kontinentalsokkelundersøkelser, Trondheim, Publication, **108**, 251–263.
- IPCC 2007. Summary for policymakers. In: SOLOMON, S., QIN, D., MANNING, M. ET AL. (eds) *Climate Change 2007: The Physical Sciences Basis, Working Group 1 Contribution to the Fourth Assessment Report of IPCC*. Cambridge University Press, Cambridge.
- JONES, G.D. 1988. A paleoecological model of late Paleocene 'Flyscht-type' agglutinated foraminifera using the paleoslope transect approach, Viking Graben, North Sea. *Abhandlungen der Geologischen Bundesanstalt*, **41**, 143–153.
- KAIHO, K., TAKEDA, K., PETRIZZO, M.R. & ZACHOS, J.C. 2006. Anomalous shifts in tropical Pacific planktonic and benthic foraminiferal test size during the Paleocene–Eocene Thermal Maximum. *Palaeogeography, Palaeoclimatology, Palaeoecology*, **237**, 456–464, doi:10.1016/j.palaeo.2005.12.017.
- MARTIN, S., RODOLFO-METALPA, R., RANSOME, E., ROWLEY, S., BUIA, M.-C., GATTUSO, J.-P. & HALL-SPENCER, J.M. 2008. Effects of naturally acidified seawater on seagrass calcareous epibionts. *Biology Letters*, **4**, 689–692, doi:10.1098/rsbl.2008.0412.
- MOY, A.D., HOWARD, W.R., BRAY, S.G. & TRULL, T.W. 2009. Reduced calcification in modern Southern Ocean planktonic foraminifera. *Nature, Geoscience*, **2**, 276–280, doi:10.1038/ngeo460.
- MURRAY, J.W. 1989. Syndeositional dissolution of calcareous foraminifera in modern shallow-water sediments. *Marine Micropaleontology*, **15**, 117–121.
- MURRAY, J.W. 1991a. Ecology and distribution of benthic foraminifera. In: LEE, J.J. & ANDERSON, O.R. (eds) *Biology of Foraminifera*. Academic Press, London, 221–253.
- MURRAY, J.W. 1991b. *Ecology and Palaeoecology of Benthic Foraminifera*. Longman, Harlow.
- MURRAY, J.W. 2006. *Ecology and Applications of Benthic Foraminifera*. Cambridge University Press, Cambridge.
- MURRAY, J.W. & ALVE, E. 1994. High diversity agglutinated foraminiferal assemblages from the NE Atlantic: dissolution experiments. *Cushman Foundation, Special Publication*, **32**, 33–51.
- MURRAY, J.W. & ALVE, E. 1999a. Taphonomic experiments on marginal marine foraminiferal assemblages: how much ecological information is preserved? *Palaeogeography, Palaeoecology, Palaeoclimatology*, **149**, 183–197.
- MURRAY, J.W. & ALVE, E. 1999b. Natural dissolution of shallow water benthic foraminifera: taphonomic effects on the palaeoecological record. *Palaeogeography, Palaeoecology, Palaeoclimatology*, **146**, 195–209.
- MURRAY, J.W. & ALVE, E. 2000. Do calcareous dominated shelf foraminiferal assemblages leave worthwhile ecological information after their dissolution? In: HART, M.B., KAMINSKI, M.A. & SMART, C.W. (eds) *Proceedings of the Fifth International Workshop on Agglutinated Foraminifera (Plymouth, UK, September 6–16, 1997)*. Grzybowski Foundation Special Publication, **7**, 311–331.
- MURRAY, J.W., ALVE, E. & CUNDY, A. 2003. The origin of modern agglutinated foraminiferal assemblages: evidence from a stratified fjord. *Estuarine, Coastal and Shelf Science*, **58**, 677–697.
- NGUYEN, T.M.P., PETRIZZO, M.R. & SPEIJER, R.P. 2009. Experimental dissolution of a fossil foraminiferal assemblage (Paleocene–Eocene Thermal Maximum, Dababiya, Egypt): implications for palaeoenvironmental reconstructions. *Marine Micropaleontology*, **73**, 241–258, doi:10.1016/j.marmicro.2009.10.005.
- ORR, J.C., FABRY, V.J., AUMONT, O., ET AL. 2005. Anthropogenic ocean acidification over the twenty-first century and its impact on calcifying organisms. *Nature*, **437**, 681–686, doi:10.1038/nature04095.
- PAGANI, M., PEDENTCHOUK, N., HUBER, M., ET AL. 2006. Arctic hydrology during global warming at the Paleocene/Eocene thermal maximum. *Nature*, **442**, 671–675, doi:10.1038/nature05043.
- RIES, J.B., COHEN, A.L. & MCCORKLE, D.C. 2009. Marine calcifiers exhibit mixed responses to CO₂-induced ocean acidification. *Geology*, **37**, 1131–1134, doi:10.1130/G30210A.1.
- RODOLFO-METALPA, R., LOMBARDI, C., COCITO, S., HALL-SPENCER, J.M. & GAMBÌ, M.C. 2010. Effects of ocean acidification and high temperatures on the bryozoan *Myriapora truncata* at natural CO₂ vents. *Marine Ecology* doi: 10.1111/j.1439-0485.2009.00354.x.
- ROYAL SOCIETY 2005. *Ocean acidification due to increasing atmospheric carbon dioxide*. Royal Society, London Policy Document, **12/05**.
- SCOTT, D., GRADSTEIN, F., SCHAFFER, C., MILLER, A. & WILLIAMSON, M. 1983. The recent as a key to the past: does it apply to agglutinated foraminiferal assemblages? In: VERDENIUS, J.G., VAN HINTE, J.E. & FORTUIN, A.R. (eds) *Proceedings of the First Workshop on Arenaceous Foraminifera, 7–9 September 1981*. Institutt for Kontinentalsokkelundersøkelser, Trondheim, Publication, **108**, 147–157.
- SPEIJER, R.P., VAN DER ZWAAN, G.J. & SCHMITZ, B., 1996. The impact of Paleocene/Eocene boundary events on middle neritic benthic foraminiferal assemblages from Egypt. *Marine Micropaleontology*, **28**, 99–132.
- SPERO, H.J., BIJMA, J., LEA, D.W. & BERNIS, B.E. 1997. Effect of seawater carbonate concentration on foraminiferal carbon and oxygen isotopes. *Nature*, **390**, 497–500.
- VERON, J.E.N., HOEGH-GULDBERG, O., LENTON, T.M., ET AL. 2009. The coral reef crisis: the critical importance of <350 ppm CO₂. *Marine Pollution Bulletin*, **58**, 1428–1436, doi:10.1016/j.marpolbul.2009.09.009.
- ZACHOS, J.C., RÖHL, U., SCHELLENBERG, S.A., ET AL. 2005. Rapid acidification of the ocean during the Paleocene–Eocene Thermal Maximum. *Science*, **308**, 1611–1615, doi:10.1126/science.1109004.